MET 3502/5561 Synoptic Meteorology

Lecture 7

AIRMASSES, FRONTS and FRONTAL ANALYSIS

Airmasses

•An airmass is a large body of air with relatively uniform thermal and moisture characteristics.

•Airmasses cover large regions of the earth, typically several hundred thousand square kilometers.

•Airmasses can be as deep as the depth of the troposphere or as shallow as 1 to 2 km.

•Airmasses form when air remains over a relatively flat region of the earth* with homogeneous surface characteristics for an extended period of time.

(* Canadian and Siberian plains, cool oceanic regions such as the North Atlantic and Pacific, deserts, such as the Sahara and the American southwest, and tropical oceanic regions including the equatorial Atlantic and Pacific, and smaller water bodies such as the Caribbean Sea and the Gulf of Mexico). 2

Sources of North American Airmasses

- **Continental Arctic** (cA): bitter cold, form over Canada and the frozen Arctic sea (north of 60N) in winter
- **Continental Polar** (cP): form over northern part of North America continent in all season.
- Maritime Polar (mP): High-latitude sea (N. Atl or N. Pac.)
- Maritime Tropical (mT): Tropical (Trade-Wind) sea (Atl. Or Pac., Gulf of Mexico, Caribean sea)
- **Continental tropical** (cT): Desert
- Air-Mass Modification takes place outside source region
- North American weather is dominated by alternation of cP and mT air



Key features of an airmass on weather maps:

The centers of <u>cold airmasses</u> are associated with high pressure on surface weather maps. High pressure develops in response to cooling.

Cooling of the whole layer \rightarrow increase air density \rightarrow pressure surface fall at upper level \rightarrow an inward pressure gradient force aloft \rightarrow upper air convergence \rightarrow air fills the column \rightarrow creating surface high pressure.



In winter, high-pressure centers form and are the dominant feature over the northern parts of the continents of Asia and North America.



Courtesy of the American Meteorological Society

Example of a high pressure system that moved southward into the central US in winter





In summer, when the oceans are cooler than the landmasses, large highpressure centers are the dominant feature of the atmosphere over the North Atlantic and Pacific Oceans.

The high-pressure center over the Atlantic is called the "Bermuda High" because it is centered near Bermuda, while its Pacific counterpart is called the "Pacific High" The centers of very warm airmasses appear as semi-permanent regions of low surface pressure.

Examples: 1) Monsoon Low in summer over S.E. Asia
2) North American Monsoon Low over Deserts of the U.S. southwest
3) Equatorial low pressure belt
4) Icelandic Low over the N. Atlantic
5) Aleutian Low over the N. Pacific in winter.

Note: Semi-permanent low-pressure centers differ substantially from migrating tropical and extratropical low-pressure centers associated with cyclones and hurricanes.

The Aleutian Low and the Icelandic Low are associated with airmasses that are warm relative to the surrounding continental airmasses in winter.



http://zubov.atmos.uiuc.edu/FINGERPRINT/obs.slp.html

Not all airmasses are in contact with the ground

Airmasses sometime exist aloft in the troposphere, residing on top of other airmasses.



Airmasses are three-dimensional, and the boundaries between airmasses are often quite sharp and distinct

Stratosphere: a large airmass that covers the entire globe



Fronts: Boundaries between airmasses

Meteorologists classify fronts based on:

- a) the thermal and moisture characteristics of the airmasses
- b) the direction of movement of the airmasses
- c) whether the boundary between the airmasses is in contact with the ground (a surface front), or can only be found aloft (an "upper level front").

Cold Front:

The boundary between two airmasses is called a cold front if the cold air is advancing forward, lifting the warm air.

The leading edge of the cold airmass typically has a shape like a dome, as shown in the figure.



A: Thunderstorms may form a squall line along the front; may form supercell thunderstorm with tonadoes, hail & damaging winds. А Warm, moist Cold air conditionally Advancing unstable air west east Location of surface front в B: Clouds; may only produce light rain or no rain at all. Cold air Warm, moist stable air Advancing west east Location С of surface front C: No Clouds. Warm, dry Cold air stable air Advancing west east Location of surface front ©2002 Kendall/Hunt Publishing 12

Identifying a cold front on a surface weather map:



- 1. Surface front is located at the leading edge of the strong temperature gradient.
- 2. Surface front is located at the leading edge of accompanying strong dewpoint gradient.
- Surface front coincides with pressure trough pressure dropping as the front approaches and rising after it passes.

Identifying a cold front on a surface weather map:



- 4. Cold air is advancing with time, replacing warm air.
- 5. A sharp wind shift, generally from winds with southerly component to winds with westerly or northwesterly component, will occur at front.
- 6. Precipitation may be behind, along, or ahead of front (or may not be present) but when present, it will be organized along a line.

Identifying a cold front on a weather map:



Leading edge of a cold front on radar often appears as a very long narrow line of high reflectivity



Cold Fronts: The leading edge of a cold front often appears like a density current, with a head, waves, and accompanying sharp discontinuity in temperature, wind, and moisture content.



Data from 600 m high tower called the Boulder Atmospheric Observatory. A cold Front passed the tower On 19 Sep 83.

Shapiro et al. 1985, MWR, Fig. 2.6 Bluestein Vol., II

Deep Tropospheric Cold Front and Jetstream



Potential temperature and isotachs

Fronts and jetstreams are related dynamically by thermal wind relationship

Keyser and Shapiro 1978 August MWR



Deep Tropospheric Cold Front

Boundary between airmasses is marked by a sharp gradient in temperature

Temperature (C)

Potential Temperature (K)

Wallace and Hobbs 1977 Fig. 3.19/3.20

"Back-Door" cold front:

A cold front that approaches an area from the east

Back door cold fronts are common in New England and along the east side of the Rocky Mountains



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Figure 2.16 Example of a New England back-door cold front 0000 UTC, April 29, 1990. Winds east of the front are off the cold Atlantic; winds west of the front are off the warmer land. Temperature and dew point (°C) plotted according to convention. Pressure plotted is altimeter setting in tens of mb without the leading "10." Full wind barb = 5 m s⁻¹; half wind barb = $2.5 m s^{-1}$.

From Bluestein² Vol II, Fig. 2.16

Warm front: Regardless of the cloud formations or precipitation, we say that a front is a warm front if the cold air is retreating and the warm air is advancing.

The leading edge of the cold airmass typically has a shape like a dome, as shown in the figure.





The type of weather along a warm front depends on the stability and moisture content of the warm air.

Widespread layered clouds will develop. These clouds are deepest just north of the frontal boundary and progressively become thinner and higher toward north. Precipitation is heaviest closer to the frontal boundary & lighter to the north.

Analyzing warm fronts



- 1. Warm front will often lie along a distinct pressure trough
- 2. Front is (normally)at southern edge of a moderate-strong N-S temperature gradient

[warm fronts are occasionally oriented NW-SE on Great Plains (i.e. an NE-SW temperature gradient)]

Analyzing warm fronts



- 1. Warm front marked by wind shift typically easterly or northeasterly winds north of the front, southerly winds south of the front.
- 2. Front is (normally) at southern edge of a moderate-strong N-S dewpoint gradient





Summary of Cold Fronts & warm Fronts

Locations:

a) Cold fronts are typically located in the southwest quadrant of a cyclone early in its lifecycle.

b) Warm fronts are usually found on the east side of a cyclone.

c) Because of the tilt of the airmass boundary, the front's horizontal position varies with elevation.

Occluded fronts A (Classical description)

- As cyclones develop, cold air to the west of the cyclone advances rapidly southward
- around the center of low pressure
- While the cold air to the north of the warm front retreats northward slowly.



The cold front will progress around the south side, then progress northeastward, approaching the warm front. When the cold air behind the cold front comes in direct contact with the cold air north of the warm front, a new airmass boundary is created called an *occluded front*.

As the fronts meet, warm air is forced aloft over both the cold air masses7

Warm & Cold Occlusion

The 3D structure of an occluded front depends on the temperatures within the cold air behind the cold front and north of the warm front.

In <u>cold occlusion</u>, the air behind the cold front is colder than air north of the warm front.



In <u>warm occlusion</u>, air north of the warm front is colder than the air behind the cold front.

Warm occlusions appear to be more common than cold occlusions. Occluded fronts develop during the mature & dissipating stages of cyclones. They are typically characterized by widespread cloudiness & rain or snowfall. 28



Classic; Traditional





Vertical Cross Section of a Lee Cyclone Showing Airmass Structure not as simple as the previous conceptual model suggests Upper level front (Cold Front Aloft) **TROWAL:** Trough Equivalent potential temperature and airmasses (color) of warm air aloft -450 primary generator of precipitation 320 Dry slot 312 500 550 310 ure (mb) 009 Cyclone Center: Pres 316 308 Intersection 700 point of 308 750· Warm front & 312 310 800 cold front 850 290 288 900 950 51 76 101 123 144 163 186 206 227 254 281 309 336 431 456 482 0 25 Distance along cross section (km) 32



New findings about occlusion structure in many cyclones forming east of the Rocky mountains

Cold Occlusion: Classical occlusion structure

Warm Occlusion: Classic occlusion structure

Upper level front

(Also called Cold Front Aloft) Common feature within cyclones that form east of the Rockies.

Boundary between:

Air descending from the upper troposphere that originates in the convergent region of the trough

and

Air ascending from the lower atmosphere in the southeastern quadrant of a cyclone







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The upper level front is "upper level" in some cyclones, but may extend to the surface as a continuous front in others. Where it reaches the surface as a single front the boundary is sometimes analyzed as a surface dry line, while other times as a surface cold front.



Dry Lines

Dry Line: A front characterized by a sharp moisture difference, but little temperature change

Dry lines develop when air flowing eastward from the high desert plateau regions of Arizona, Colorado, and New Mexico descends the Rockies into the southern plains and encounters moist air flowing northward from the Gulf of Mexico



The Great Plains Dryline



Westerly winds (often strong)

Clear skies

Daytime: Warm Temps

Large diurnal temperature range

Low moisture



Ahead of the dry line



Fig. 2.40 Bluestein II

Severe thunderstorms often develop along dry lines in late afternoon, particularly when dry air is present aloft above the moist air east of the dry line.



Courtesy of Bruce Lee

Stationary Fronts

Fronts are sometimes stationary. Although the boundary is stationary, air on both sides of the boundary can be moving.

With a *stationary front*, air on the cold side of the front will always be flowing parallel to the front.

Regardless of the cloud formations or precipitation, we say that the front is **a stationary front** if the cold air is neither advancing or retreating.

Warm air conditionally unstable: A line of showers & thunderstorms May develop in the warm air flash flooding.

Warm air stable: widespread layered clouds may form over the front, with rain falling on the cold side.



Stationary fronts may have physical structure similar to cold fronts or warm fronts

Significant weather along Α stationary fronts **FLASH FLOODS** Warm, moist conditionally Cold air flowing out of unstable air paper, parallel to front East West Location of surface front В **ICE STORMS** Warm, moist conditionally unstable air Cold air flowing into paper, parallel to front North South Location of surface front 41 ©2002 Kendall/Hunt Publishing

Example of a stationary front that produced a local flash flood in Kentucky in 1989





Stationary front that led to 30 Million Dollar Ice storm in 1990

Topography and Fronts

Cold air dammed along the Colorado Rockies is often analyzed as a stationary front. It is a boundary between air and rock.



Courtesy of the Department of Atmospheric Sciences University of Illinois at Urbana-Champaign

Topography and Fronts

Cold air trapped in Appalachian valleys is often analyzed as a distorted warm front. The cold air in the front's southward "dip" is quite shallow. Above the top of the shallow air the front does not dip southward.



Courtesy of the Department of Atmospheric Sciences University of Illinois at Urbana-Champaign